

# Wave generation techniques using the Boussinesq-Abbott & Green-Naghdi equations

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# Dimensionless Nonlinear Boussinesq Equations

- Dimensionless Boussinesq system

$$\begin{cases} \partial_t \zeta + \partial_x q = 0, \\ \left(1 - \frac{\mu}{3} \partial_x^2\right) \partial_t q + \partial_x \left(\frac{1}{2\varepsilon} h^2 + \varepsilon \frac{1}{h} q^2\right) = 0, \end{cases}$$

- Coupling surface elevation  $\zeta$  above rest state to horizontal discharge  $q$
- $\varepsilon$  and  $\mu$  respectively called *nonlinearity* and *shallowness* parameters
- Initial conditions

$$(\zeta, q)(t = 0, x) = (\zeta^0, q^0)(x)$$

⇒ How to impose the generating boundary condition  $\zeta(t, x = 0) = f(t)$  ?

# Propagation of solitary wave

Propagation of the soliton,  $N_x = 200$

## Generating boundary condition ?

$$\begin{cases} \partial_t \zeta + \partial_x q = 0, \\ \left(1 - \frac{\mu}{3} \partial_x^2\right) \partial_t q + \partial_x \left(\frac{1}{2\varepsilon} h^2 + \varepsilon \frac{1}{h} q^2\right) = 0, \end{cases}$$

with initial and boundary conditions

$$(\zeta, q)(t = 0, x) = (\zeta^0, q^0)(x), \quad \zeta(t, x = 0) = f(t).$$

- For periodic boundary conditions : inversion of operator  $(1 - \frac{\mu}{3} \partial_x^2)$  ok
- For a generating boundary condition : not possible to use Riemann invariants as in Nonlinear Shallow Water case
- Inversion of  $(1 - \frac{\mu}{3} \partial_x^2)$  on the half-line  $(0, \infty)$  : **need of a boundary condition on  $\partial_t q$ !**

## Generating boundary condition

$$\begin{cases} \partial_t \zeta + \partial_x q = 0, \\ (1 - \frac{\mu}{3} \partial_x^2) \partial_t q + \partial_x \left( \frac{1}{2\varepsilon} h^2 + \varepsilon \frac{1}{h} q^2 \right) = 0, \end{cases}$$

with initial and boundary conditions

$$(\zeta, q)(t = 0, x) = (\zeta^0, q^0)(x), \quad \zeta(t, x = 0) = f(t).$$

Strategy :

- use the **inverse of operator**  $(1 - \frac{\mu}{3} \partial_x^2)$  with homogeneous Dirichlet BC
- construct the **dispersive boundary layer** accounting for the non-zero boundary value of  $q$

Reformulation of the Boussinesq equations as a system

- with a **nonlocal flux**
- and a **source term** accounting for the dispersive boundary layer

# Nonlinear Boussinesq Equations

## Definition

We denote by  $R_0$  and  $R_1$  the inverse of the operator  $(1 - \frac{\mu}{3}\partial_x^2)$  with homogeneous Dirichlet and Neumann boundary conditions :

$$R_0 : \begin{array}{l} L^2(\mathbb{R}_+) \rightarrow H^2(\mathbb{R}_+) \\ f \mapsto u, \end{array} \quad \text{where} \quad \begin{cases} (1 - \frac{\mu}{3}\partial_x^2)u = f \\ u(0) = 0 \end{cases}$$

and

$$R_1 : \begin{array}{l} L^2(\mathbb{R}_+) \rightarrow H^2(\mathbb{R}_+) \\ f \mapsto v, \end{array} \quad \text{where} \quad \begin{cases} (1 - \frac{\mu}{3}\partial_x^2)v = f \\ \partial_x v(0) = 0 \end{cases} .$$

## Lemma

For all  $f \in L^2(\mathbb{R}_+)$ ,

$$R_0 \partial_x f = \partial_x R_1 f.$$

**Notation :**  $R_1 f = (R_1 f)|_{x=0}$  and  $q(t) = q(t, x = 0)$

## Boussinesq Equations : dispersive boundary layer

$$\begin{cases} \partial_t \zeta + \partial_x q = 0, \\ (1 - \frac{\mu}{3} \partial_x^2) \partial_t q + \partial_x (\tilde{f}(\zeta, q)) = 0, \end{cases} \quad \text{with } \tilde{f}(\zeta, q) = \frac{1}{2\varepsilon} h^2 + \varepsilon \frac{1}{h} q^2$$

- The ODE

$$Y - \frac{\mu}{3} Y'' = f, \quad Y(0) = Y_0$$

admits a unique solution given by

$$Y(x) = R_0 f + Y_0 \exp\left(-\frac{x}{\delta}\right) \quad \text{with} \quad \delta = \sqrt{\frac{\mu}{3}},$$

- Thus the second equation of Boussinesq system can be written equivalently

$$\partial_t q = -R_0 \partial_x (\tilde{f}(\zeta, q)) + \underbrace{\dot{q} \exp\left(-\frac{x}{\delta}\right)}_{\text{dispersive boundary layer}}.$$

⇒ Need to compute  $\dot{q}$

## How to compute $\underline{\dot{q}}$ ?

$$\partial_t \underline{q} + \partial_x R_1(\underline{f}(\underline{\zeta}, \underline{q})) = \underline{\dot{q}} \exp\left(-\frac{x}{\delta}\right).$$

- Differentiating with respect to  $x$

$$\partial_t \partial_x \underline{q} + \partial_x^2 R_1(\underline{f}(\underline{\zeta}, \underline{q})) = -\frac{1}{\delta} \underline{\dot{q}} \exp\left(-\frac{x}{\delta}\right).$$

- Substitution  $\partial_t \partial_x \underline{q} = -\partial_t^2 \underline{\zeta}$

$$-\partial_t^2 \underline{\zeta} + \partial_x^2 R_1(\underline{f}(\underline{\zeta}, \underline{q})) = -\frac{1}{\delta} \underline{\dot{q}} \exp\left(-\frac{x}{\delta}\right).$$

- Using  $(1 - \frac{\mu}{3} \partial_x^2) R_1 = \text{Id}$

$$\partial_t^2 \underline{\zeta} = \frac{1}{\delta^2} (R_1 - \text{Id})(\underline{f}(\underline{\zeta}, \underline{q})) + \frac{1}{\delta} \underline{\dot{q}} \exp\left(-\frac{x}{\delta}\right).$$

- At  $x = 0$ , formula for  $\underline{\dot{q}}$  :

$$\underline{\ddot{\zeta}} = \frac{1}{\delta^2} \left[ \underline{R}_1(\underline{f}(\underline{\zeta}, \underline{q})) - \underline{f}(\underline{\zeta}, \underline{q}) \right] + \frac{1}{\delta} \underline{\dot{q}}$$



# Equivalent form of Boussinesq Equations

$$\begin{cases} \partial_t \zeta + \partial_x q = 0, \\ \partial_t q + \partial_x R_1 \bar{f}(\zeta, q) = \underline{Q}(q, f, \ddot{f}, \zeta, q) \exp\left(-\frac{x}{\delta}\right), \\ \underline{\dot{q}} = \underline{Q}(q, f, \ddot{f}, \zeta, q), \end{cases}$$

with

$$\bar{f}(\zeta, q) = \frac{1}{2\varepsilon}(h^2 - 1) + \varepsilon \frac{1}{h} q^2 \quad (= \text{flux for NLSW momentum})$$

$$\underline{Q}(q, f, \ddot{f}, \zeta, q) = \frac{\varepsilon}{\delta} \frac{q^2}{1 + \varepsilon f} + \delta \ddot{f} + \frac{1}{\delta} \left(1 + \frac{\varepsilon}{2} f\right) f - \frac{1}{\delta} R_1 \bar{f}(\zeta, q),$$

+ initial data and and boundary condition  $\zeta(t, x = 0) = f(t)$ .

## Discretization

- Lax-Friedrichs scheme for  $U_i^n$ , explicit Euler scheme for the ODE on  $\underline{q}$ ,
- Non-local flux  $\tilde{f}_\mu(U) = R_1 \tilde{f}(U)$  computed with second-order centered finite-differences

$$\begin{cases} \frac{U_i^{n+1} - U_i^n}{\delta_t} + \frac{1}{\delta_x} (\tilde{\mathfrak{F}}_{\mu,i+1/2}^n - \tilde{\mathfrak{F}}_{\mu,i-1/2}^n) = \underline{S}_i^n, & i \geq 1, \quad n \geq 1, \\ \frac{\underline{q}^{n+1} - \underline{q}^n}{\delta_t} = \underline{Q}(\underline{q}^n, f^n, \ddot{f}^n, \zeta^n, q^n) & n \geq 1, \end{cases}$$

where  $U^n = (\zeta_i^n, q_i^n)_{i \geq 1}^T$ ,  $\tilde{\mathfrak{F}}_\mu(U) = (q, \tilde{f}_\mu(U))^T$  and

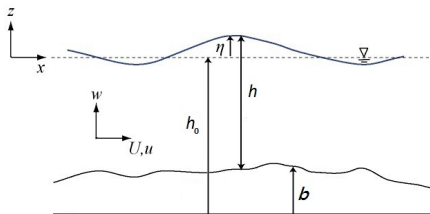
$$\underline{S}_i^n = \begin{pmatrix} 0 \\ \underline{Q}(\underline{q}^n, f^n, \ddot{f}^n, \zeta^n, q^n) \exp(-\frac{x_i}{\delta}) \end{pmatrix}$$

- Computational cost similar to the case of periodic conditions

# Propagation of solitary wave

Decomposition of  $\partial_t q$  into  $\partial_t q = -\partial_x R_1(\tilde{f}(\zeta, q)) + \dot{q} \exp\left(-\frac{x}{\delta}\right)$  (dispersive boundary layer and solution of homogenous problem)

# The physical problem



$\eta$  : **free surface elevation** ;

$h_0$  : **steel water level** ;

$b(x)$  : **bottom's topography variation** ;

$h(x, t) = h_0 + \eta(x, t) - b(x)$  : **total water depth** ;

$u(x, t)$  : **flow velocity** ;

## Mathematical model : The SW & GN equations

One dimensional form :

$$\begin{aligned} h_t + (hu)_x &= 0 \\ (I + \alpha\mathcal{T}) \left[ (hu)_t + (hu^2)_x + g \frac{\alpha - 1}{\alpha} h\eta_x \right] + \frac{g}{\alpha} h\eta_x + h\tilde{Q}(u) &= 0 \end{aligned}$$

where

$$\begin{aligned} \mathcal{T}(\cdot) &= -\frac{1}{3}h^2(\cdot)_{xx} - \frac{1}{3}hh_x(\cdot)_x + \frac{1}{3}[h_x^2 + hh_{xx}](\cdot) + \left[ b_x h_x + \frac{1}{2}hb_{xx} + b_x^2 \right](\cdot) \\ \tilde{Q}(\cdot) &= 2hh_x(\cdot)_x^2 + \frac{4}{3}h^2(\cdot)_x(\cdot)_{xx} + b_x h(\cdot)_x^2 + b_{xx}h(\cdot)(\cdot)_x \\ &+ \left[ b_{xx}h_x + \frac{1}{2}hb_{xxx} + b_x b_{xx} \right](\cdot)^2 \end{aligned}$$

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$$\begin{aligned} \tilde{Q}(\cdot) &= 2h h_x(\cdot)_x^2 + \frac{4}{3} h^2(\cdot)_x(\cdot)_{xx} + b_x h(\cdot)_x^2 + b_{xx} h(\cdot)(\cdot)_x \\ &+ \left[ b_{xx} h_x + \frac{1}{2} h b_{xxx} + b_x b_{xx} \right] (\cdot)^2 \end{aligned}$$

Dispersion relation

$$\omega^2 = g h_0 k^2 \frac{1 + \frac{\alpha-1}{3} k^2 h_0^2}{1 + \frac{\alpha}{3} k^2 h_0^2}.$$

where  $\alpha = 1.159$

## Mathematical model : Elliptic-hyperbolic decoupling

Re-write the system as :

$$\begin{aligned} h_t + (hu)_x &= 0 \\ (I + \alpha\mathcal{T}) \left[ (hu)_t + (hu^2)_x + gh\eta_x \right] - \mathcal{T}(gh\eta_x) + h\tilde{Q}(u) &= 0 \end{aligned}$$

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$$(I + \alpha\mathcal{T})\phi = \mathcal{W} - \mathcal{R}$$

$$h_t + (hu)_x = 0$$

$$(hu)_t + (hu^2)_x + gh\eta_x = \phi$$

where  $\mathcal{W} = g\mathcal{T}(h\eta_x)$  and  $\mathcal{R} = h\tilde{Q}(u)$ .



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### Two step solution

- 1 An elliptic step solving for the non-hydrostatic term  $\phi$ ;
- 2 An hyperbolic step evolving the flow variables .

## Structure of the code :

Go to : <https://mskazolea.wixsite.com/personal>  
files : main.m, flux.m, get\_PHI.m, limiter.m, muscl.m,  
primitive.m, boundary.m wave\_generation.m

## Structure of the code :

Go to : [https://mskazolea.wixsite.com/personal\\_files](https://mskazolea.wixsite.com/personal_files) : main.m, flux.m, get\_PHI.m, limiter.m, muscl.m, primitive.m, boundary.m wave\_generation.m

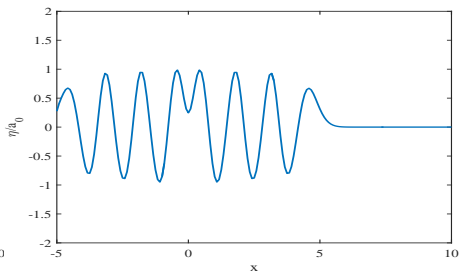
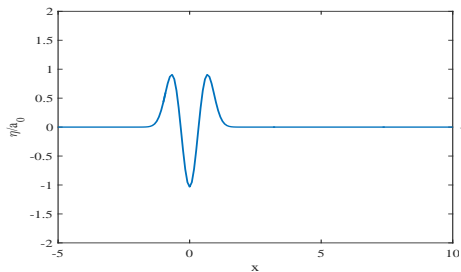
### main.m :

- Input values provided by the users
- Initialization of the solution
- RK3 loop in time
  - ▶ For each RK step  $p = 1, 2, 3$

{ Find the fluxes  $\rightarrow$  *flux.m* (Third order FV scheme)  
Update the value of  $h^p$   
Find the non Hydrostatic terms (if GN are solved)  $\rightarrow$  *get\_PHI.m* ( $C^0$  Galerkin)  
Update the value of  $(hu)^p$   
Sponge layers  $\rightarrow$  *sponge.m*

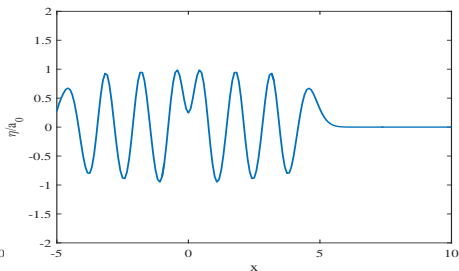
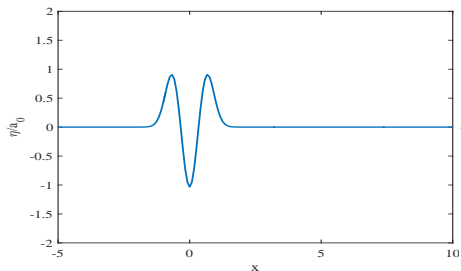
- ▶ Find new dt

# Wave generation : internal wave generation



- Common technique for BT models. (Wei & Kirby 1999)
- A source term added to the governing equations
  - 1 In a form of a mass source
  - 2 An applied pressure forcing the momentum equations

# Wave generation : internal wave generation



- Common technique for BT models. (Wei & Kirby 1999)
- A source term added to the governing equations
  - 1 In a form of a mass source
  - 2 An applied pressure forcing the momentum equations
- Derivation from the linearized Boussinesq.
- Different tuning for different BT model.

# Wave generation : internal wave generation

$$h_t + (hu)_x = f(x, t)$$

## Wave generation : internal wave generation

$$h_t + (hu)_x = f(x, t)$$

where

$$f(x, t) = D \sin(-\omega t) \exp(Bs(x - x_0)^2)$$

for a given wave frequency  $\omega$  wave direction  $\theta$ , wave amplitude  $\eta_0$  water depth  $h$  and source width parameter  $Bs$  the corresponding source amplitude  $D$  can be determined by

$$D = \frac{2\eta_0(\omega^2 - \alpha_1 g k^4 h^3) \cos\theta}{\omega l_1 k [1 - \alpha(kh)^2]}$$

## Absorbing boundaries : sponge layer approach

- They should dissipate the energy of incoming waves perfectly, in order to eliminate unphysical reflections.

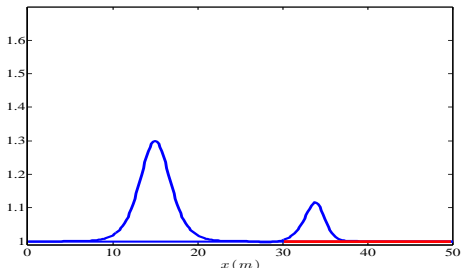


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- A sponge layer is defined in-front of the physical boundary. First proposed by Larsen and Dancy 1983

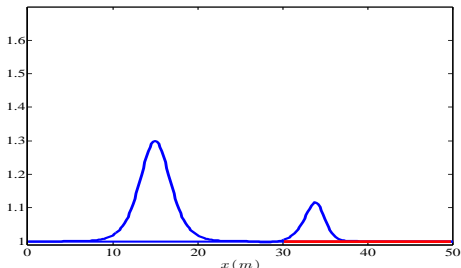
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- On it the surface elevation and/or the fluxes are damped multiplying their values by a coefficient  $\mu(x)$  after the computation of  $q^{n+1}$
- Choice of  $\mu(x)$  ?

## Absorbing boundaries : sponge layer approach (cont)

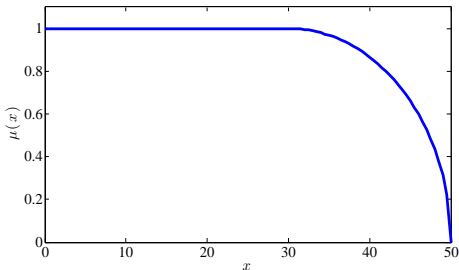
- Tonelli and Petti (2009) :

$$\mu(x) = \begin{cases} 0.5 + 0.5 * \cos\left(\pi \frac{L_s}{(L_s - D(x))}\right), & D \leq L_s \\ 1, & D \geq L_s \end{cases}$$

## Absorbing boundaries : sponge layer approach (cont)

- Wu (2004) :

$$\mu(x) = \sqrt{1 - \left(\frac{1 - D(x)}{L_s}\right)^2}$$

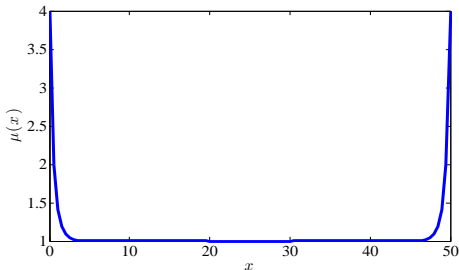


$L_s$ -sponge layer width,  $D(x)$ =normal distance between the cell center with coordinates  $x$  and the absorbing boundary.

## Absorbing boundaries : sponge layer approach (cont)

- Larsen and Dancy (1983)

$$\mu(x) = \begin{cases} \exp\left(\left(2^{d/\Delta d} - 2^{-L_s/\Delta d}\right) \ln a\right), & 0 \leq d \leq L_s \\ 1, & d_s < d \end{cases}$$



$d$ - distance between the point on the sponge layer and the boundary,  $\Delta d$   
-dimension of the elements,  $a$ -parameter

# Absorbing boundaries using extra damping terms

- Artificial damping terms  $F_b$  are added to the right hand side of the equation(s).
- General notation (Israeli and Orzag) :

$$F_b = -\omega_1(x)u + \omega_2(x)u_{xx} + \omega_3(x)\sqrt{\frac{g}{h}}\eta \quad (1)$$

where  $\omega_i = c_i\omega f(x)$

$$f(x) = \frac{\exp((x - x_s)/L_s)^2 - 1}{\exp(1) - 1}$$

## Absorbing boundaries using extra damping terms (cont)

- Different works use different coefficients e.g.
  - ▶ Kirby et al. use (1) mainly with  $c_1 = 10$ ,  $c_2 = c_3 = 0$ ; adding the terms only to the momentum equation.
  - ▶ Klonaris et al. add (1) only in the momentum equation with  $c_3 = 0$
  - ▶ Filippini et al  $\omega_1 = \omega_3 = 0$  and  $\omega_2 = 0.1$  with

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adding the terms to all the equations. c-parameter.



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- **Attention !** on how these terms affect the stability of the scheme

## References :

- 1 Israeli and Orzag, 1981. Approximation of radiation boundary conditions, J. Comp. Phys,41,15-135.
- 2 Kirby,J., Wei, G., Chen, Q., Kennedy, A. and Dalrymple, A. Funwave 1.0, 1998. Fully non-linear wave Model, Documentation and User's Manual.
- 3 A.G.Filippini, M. Kazolea and M. Ricchiuto, A flexible and genuinely nonlinear approach for nonlinear wave propagation , breaking and run-up, J.C.P., 130, 381-417,2016
- 4 Tonelli, M and Petti, M, 2009. Hybrid finite volume-finite difference scheme for 2HD improved Boussiesq equarions, Coastal Eng.,56.
- 5 Larsen, J. and Dancy, 1983.H. Open boundaries in short wave simulation- a new approach. Coastal Eng, 7, 285-297.
- 6 Wu, T.-R., 2004. A numerical study of three dimensional breaking waves and turbulence effects, Ph.D. thesis. Cornell University.